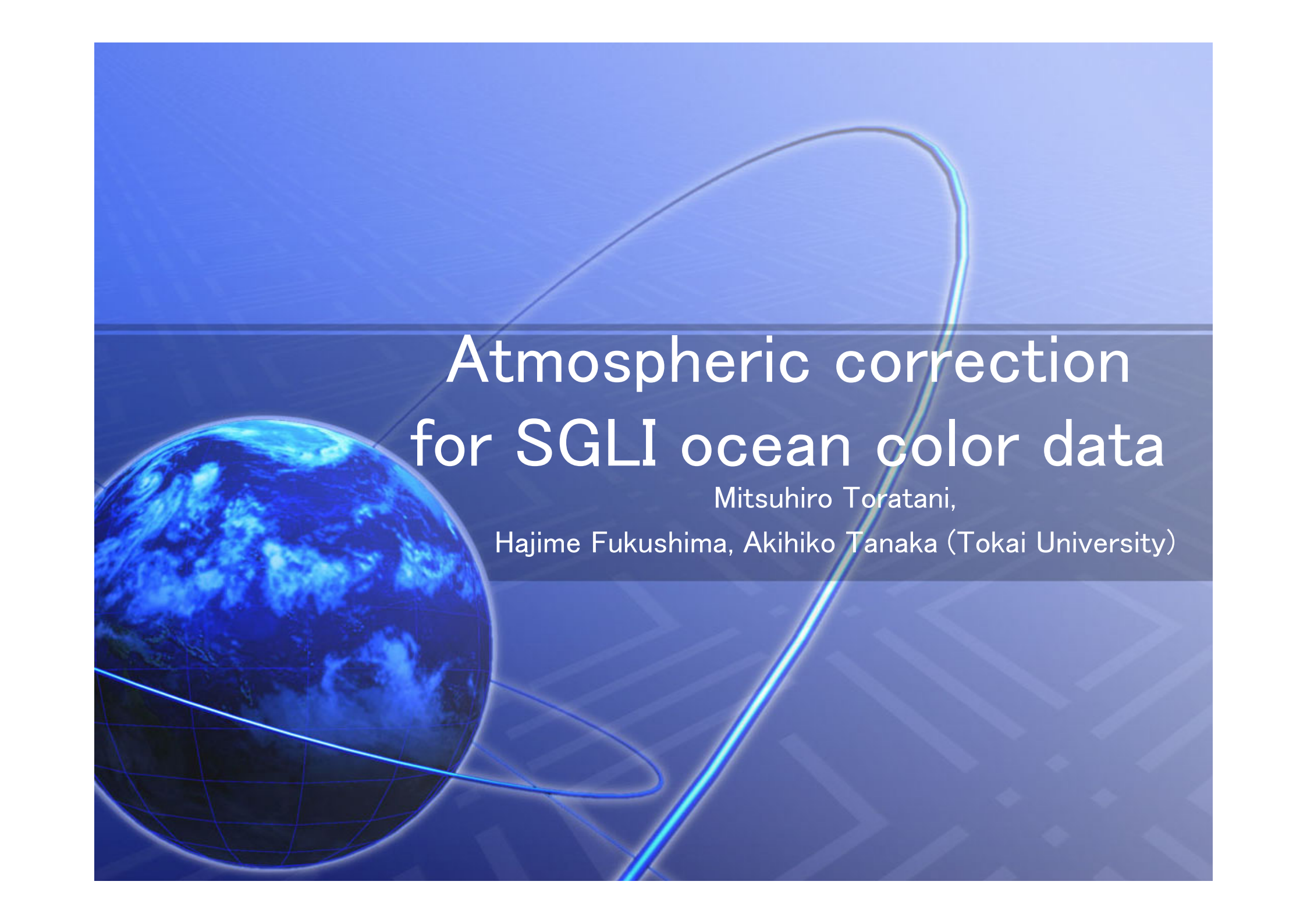




Atmospheric correction for SGLI ocean color data

Mitsuhiro Toratani,
Hajime Fukushima, Akihiko Tanaka (Tokai University)

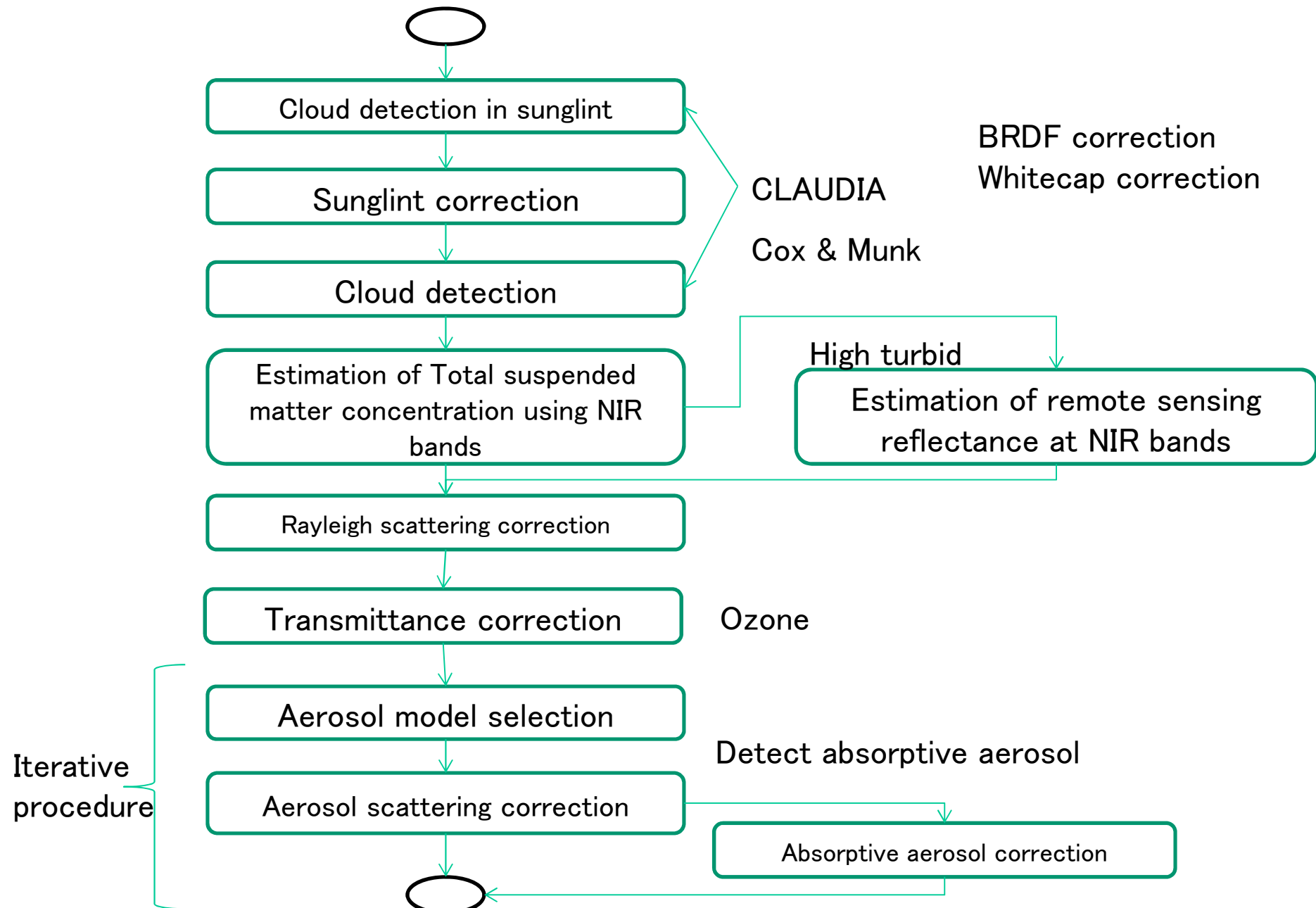
Contents

- Overview of atmospheric correction
- Spectral dependency of aerosol optical thickness
 - Relationship between Angstrom exponent aerosol type
- Aerosol model to make look-up tables for aerosol reflectance correction
 - Proposal model and sensitivity test of aerosol parameters
- Cloud detection in cloud edge
 - Performance of Inter-band Radiance Combination index method (IRC)

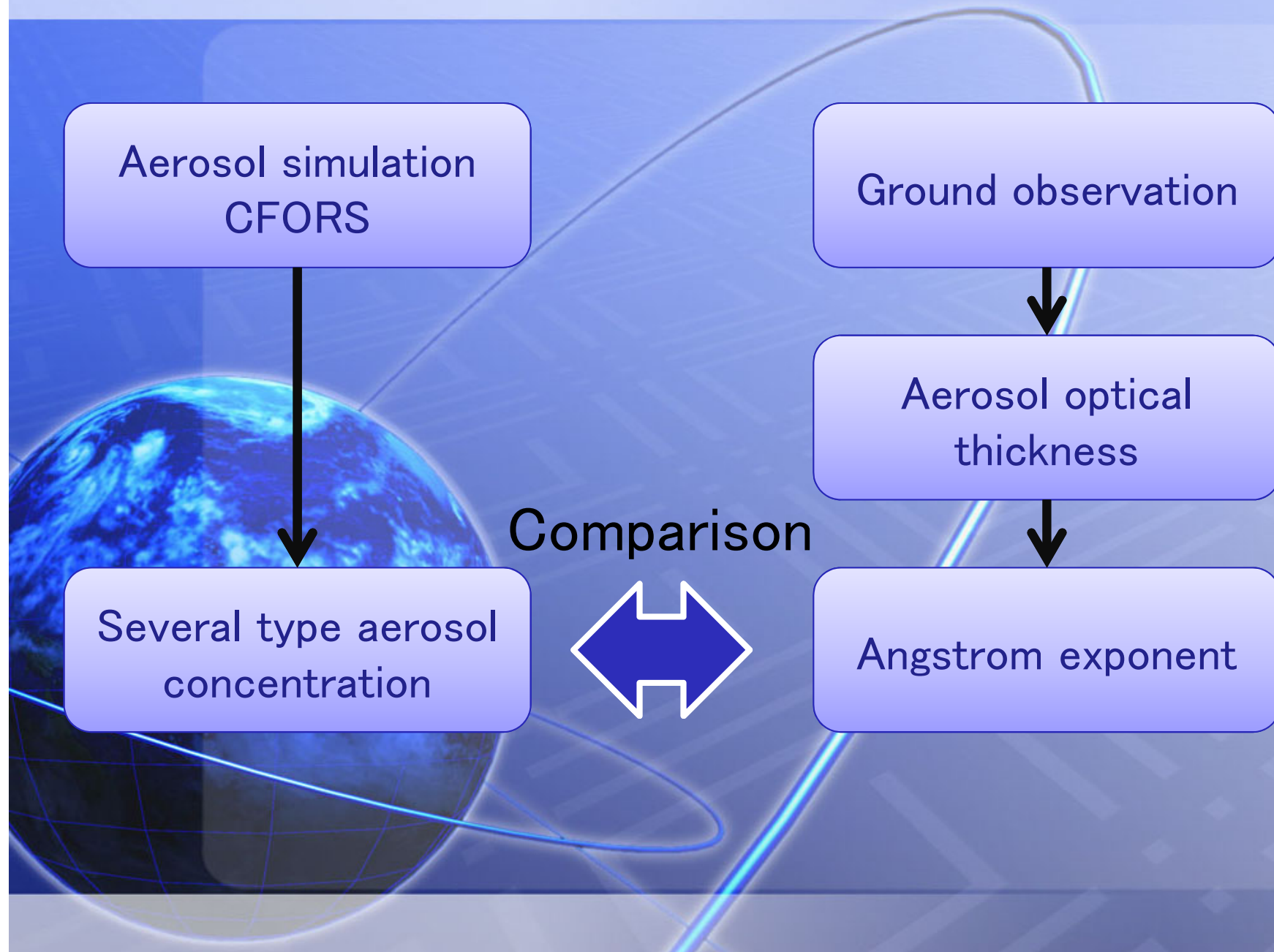
Overview of atmospheric correction for SGLI

- Based on
 - Wang and Gordon (1994) for SeaWiFS
 - Siegel et al.(2000)
 - Stumpf et al.(2003)
 - Toratani et al.(2007) for GLI
- Correction of absorptive aerosol
- Consideration of high turbid area

Flowchart of atmospheric correction scheme

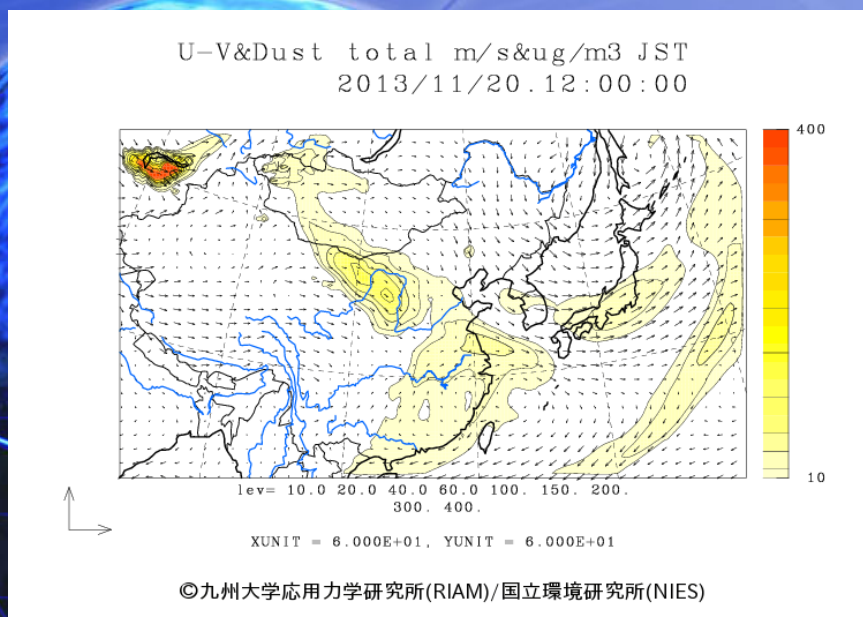


Comparison simulated-aerosol with ground observation



RIAM-CFORS

- Chemical weather FORecasting System
- based on 3D on-line regional scale chemical transport model fully coupled with RAMS (Regional Atmospheric Modeling System)



DUST

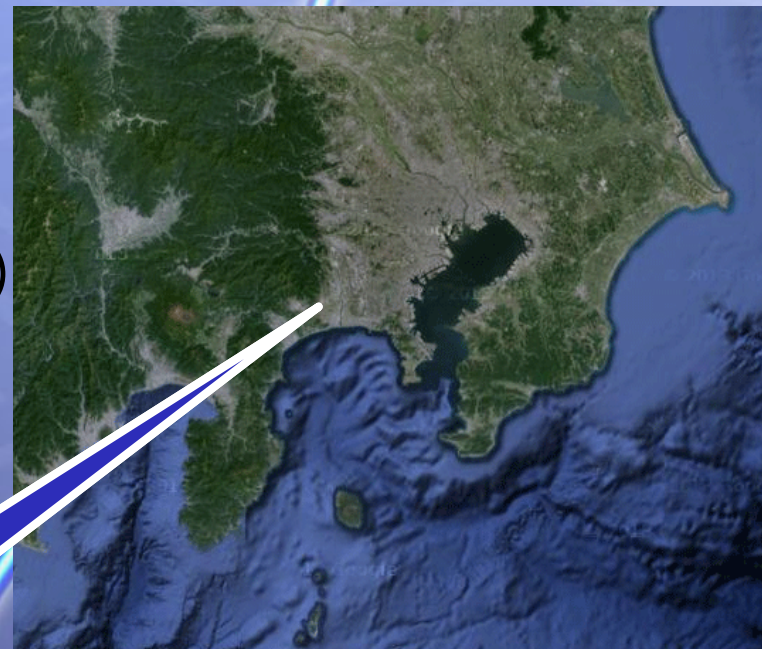
Sulfate : SO_4

Black Carbon: BC

Organic Carbon : OC

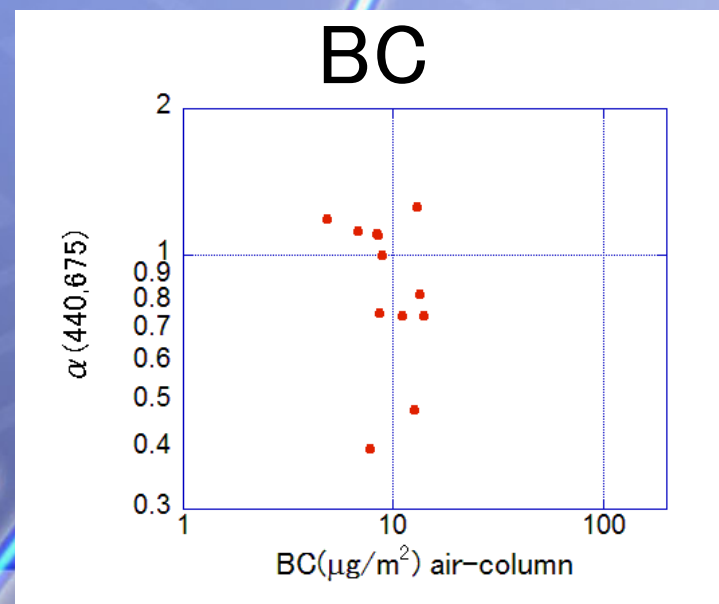
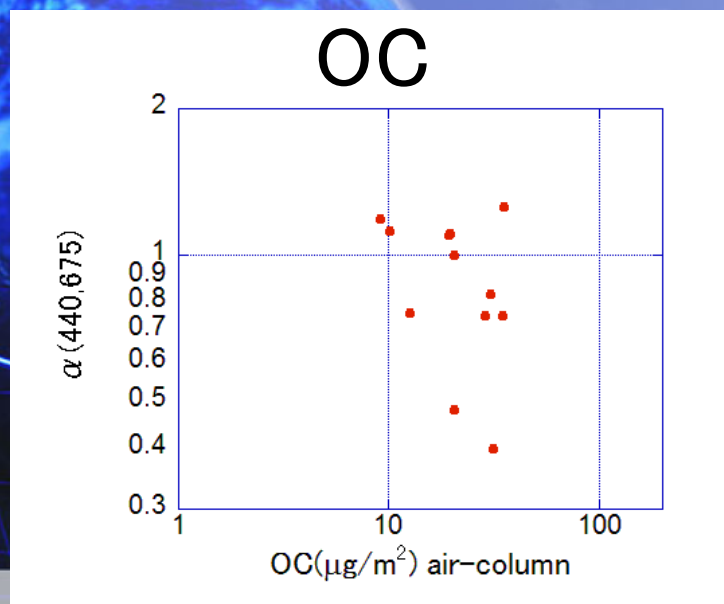
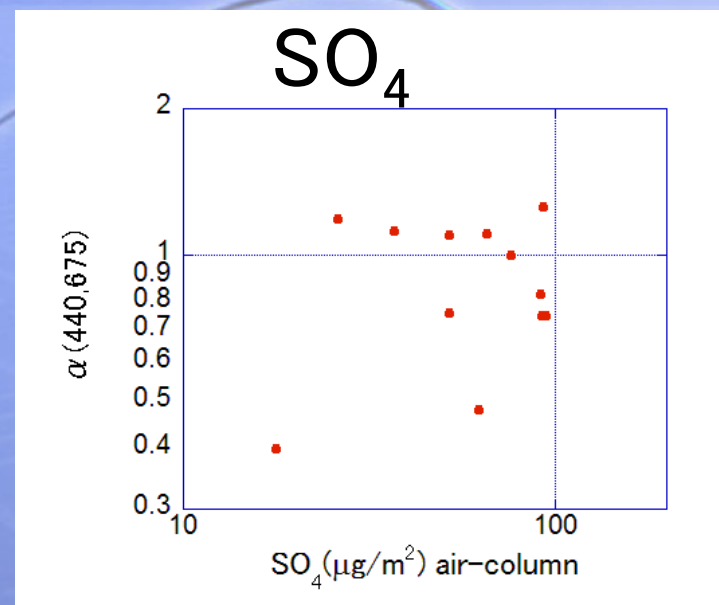
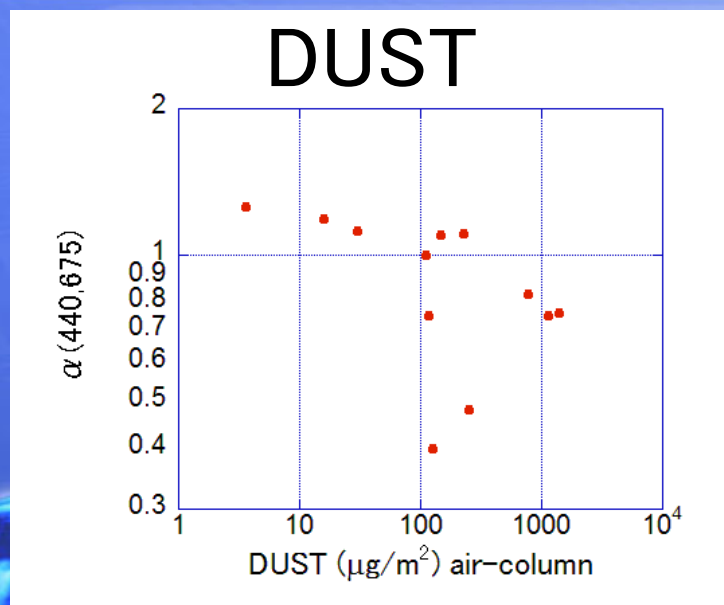
Microtops 2 Observation

- 2013/2/26–7/8
- 12 fine days
 - Dust : 7 days
- Wavelength
 - 380, 440, 675,
 - 870, 1020nm
- Calculate $\alpha(440, 675)$

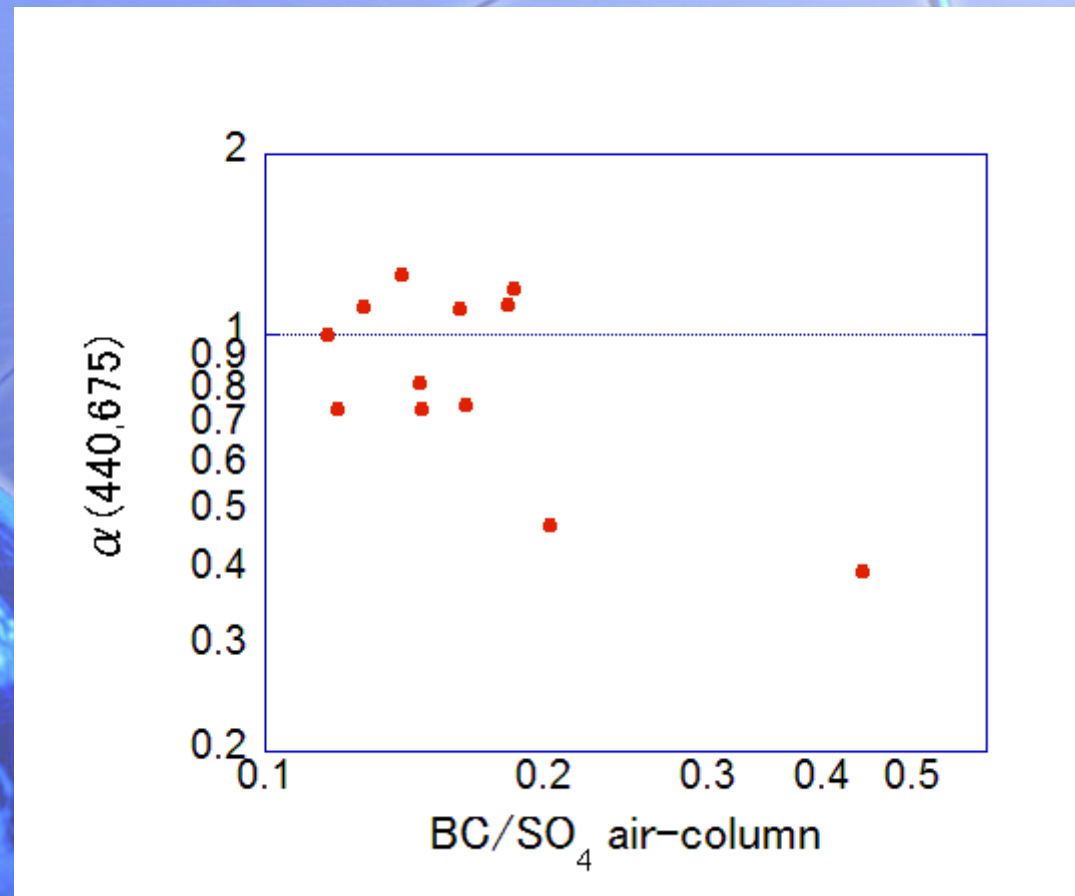


Tokai
University

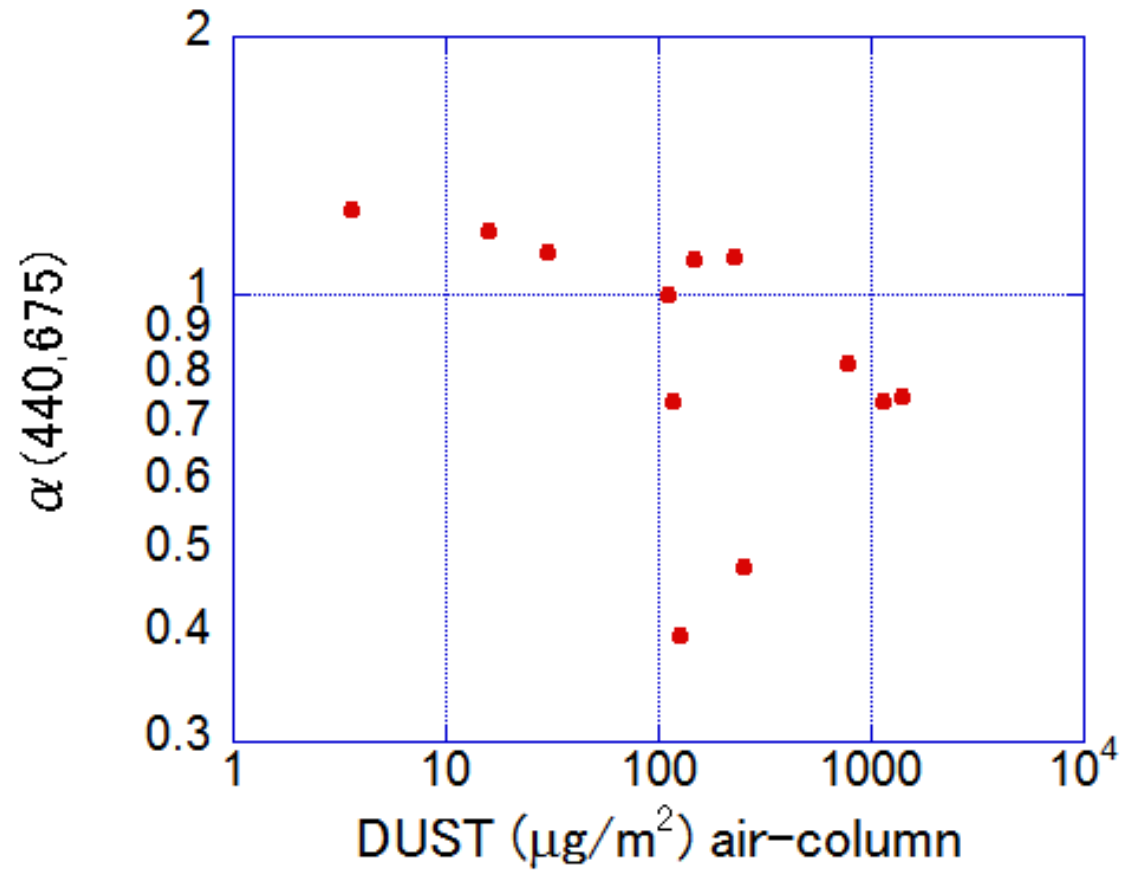
Angstrom exponent vs Aerosol concentration



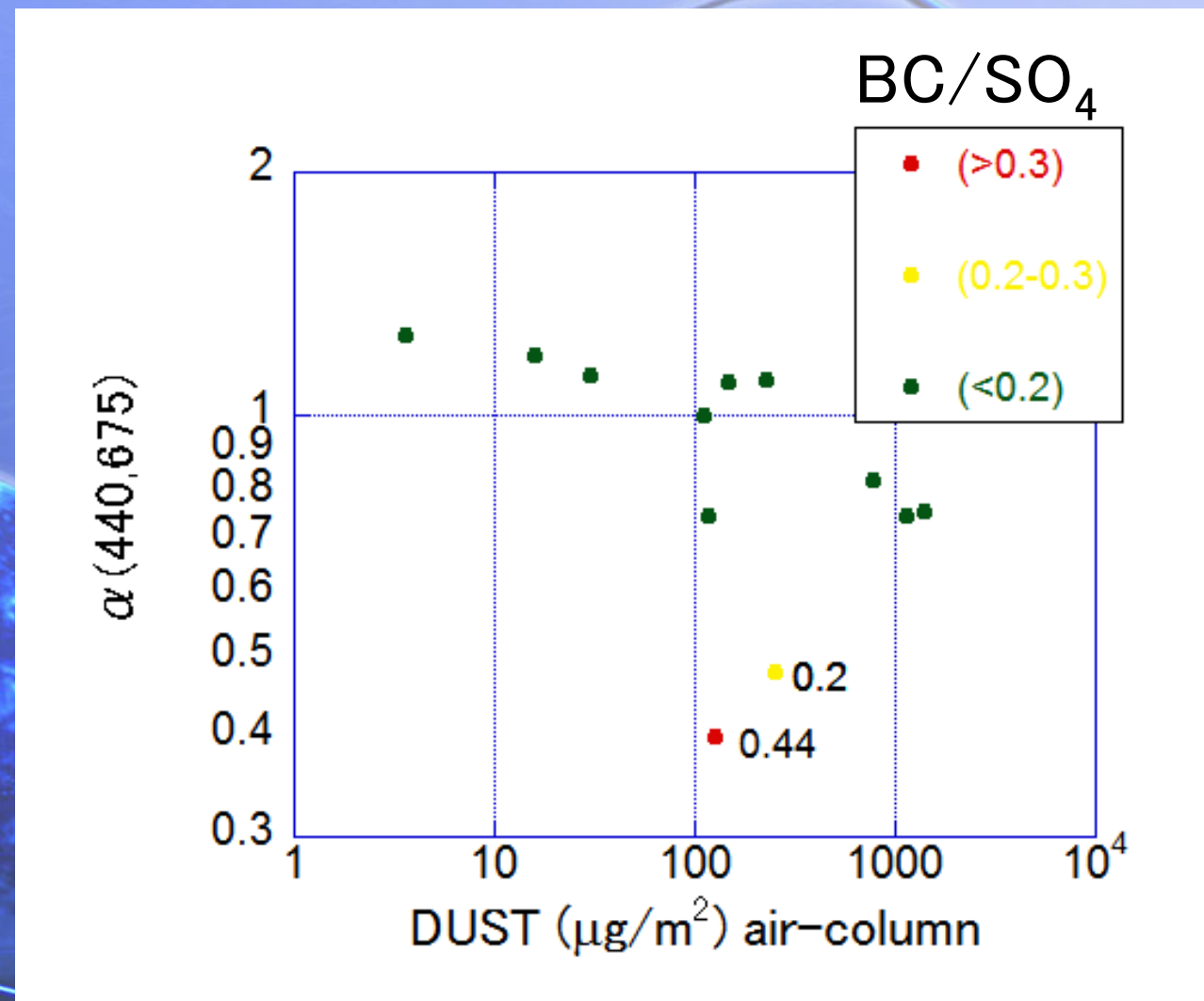
Relationship between α and BC/SO_4



Relationship between α and DUST

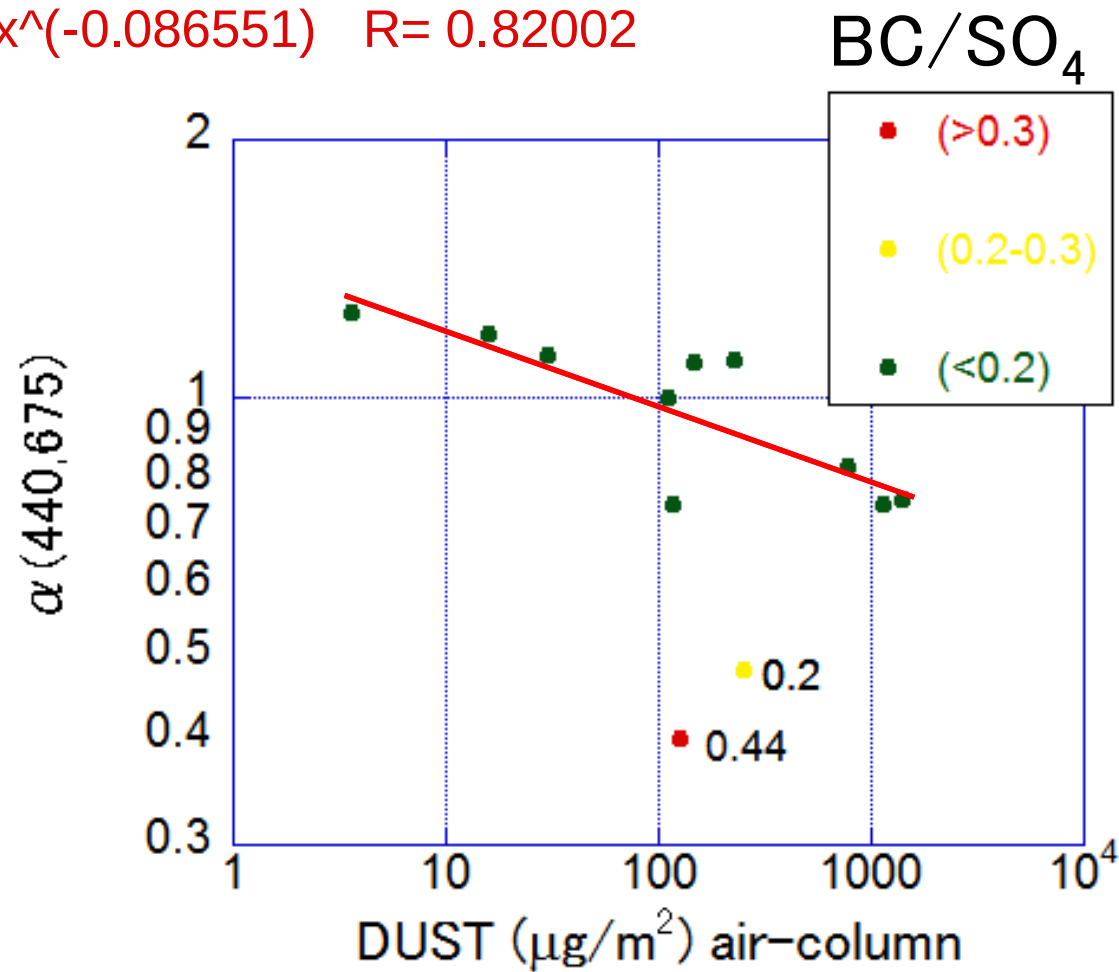


Relationship between α and DUST



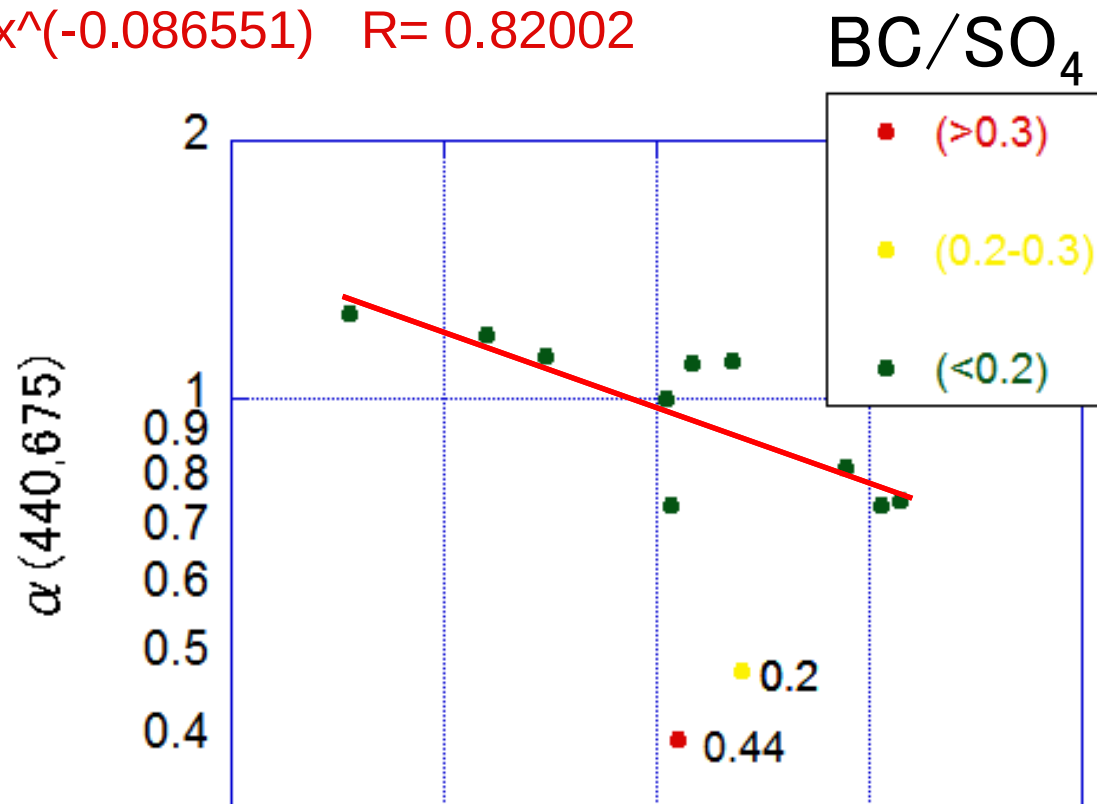
Relationship between α and DUST

— $y = 1.4709 * x^{(-0.086551)}$ $R = 0.82002$



Relationship between α and DUST

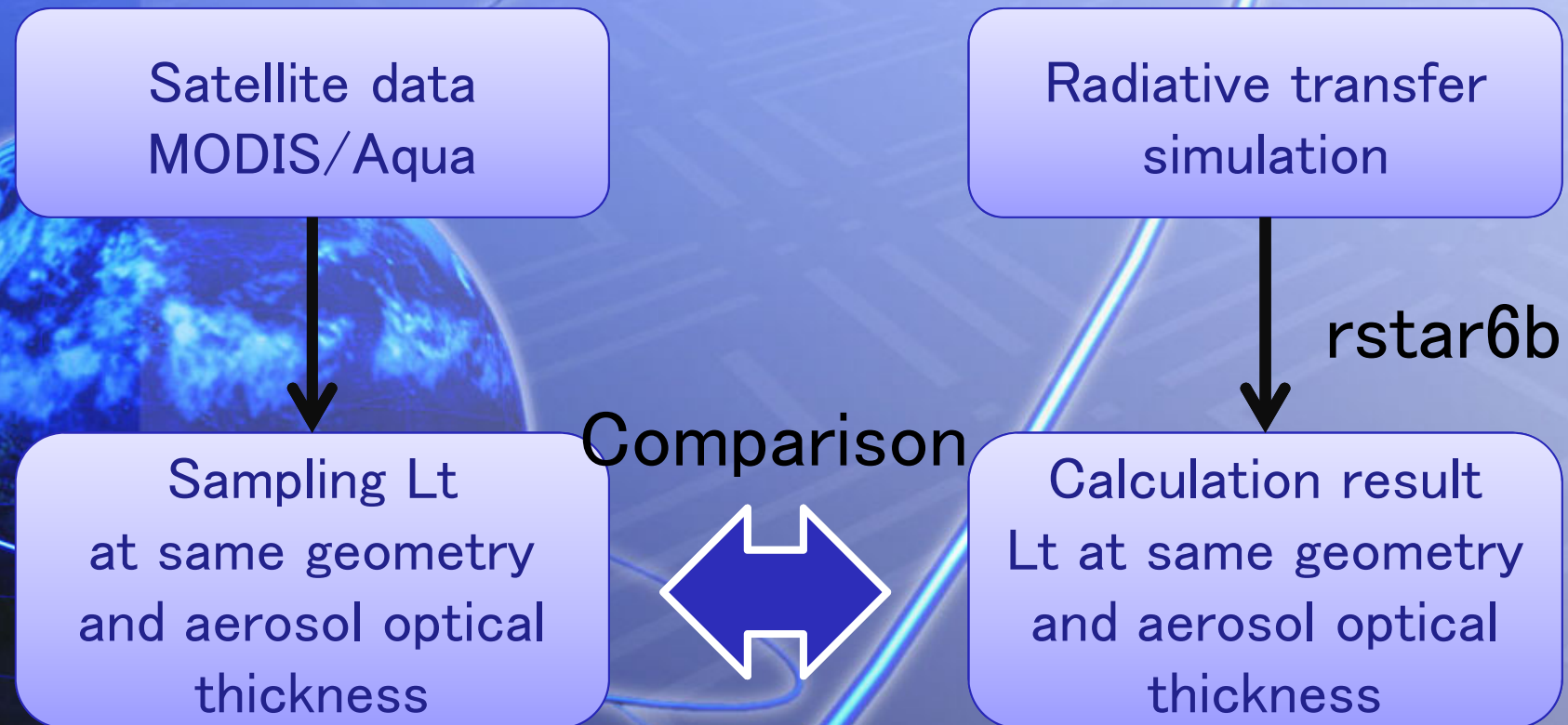
$$y = 1.4709 * x^{(-0.086551)} \quad R = 0.82002$$



Dust quantity related $\alpha(440,675)$ except for the ratio BC/SO₄ was higher (>0.2). Requiring attention of component of fine mode, especially absorptive aerosol such as black carbon.

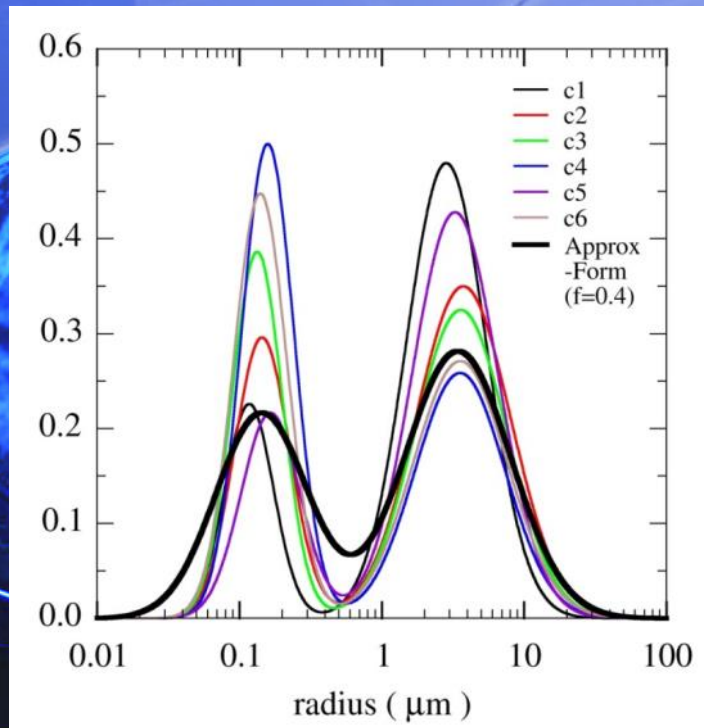
Comparison MODIS data with Radiative transfer simulation

- We need to fix aerosol model to make lookup tables for aerosol reflectance and transmittance correction



Aerosol model

- Bimodal model
 - Aerosol model provided by atmospheric group
 - It derived from AERONET data



Proposed model

Fine mode

$$r=0.14 \quad \sigma=1.86$$

Coarse mode

$$r=3.42 \quad \sigma=2.34$$

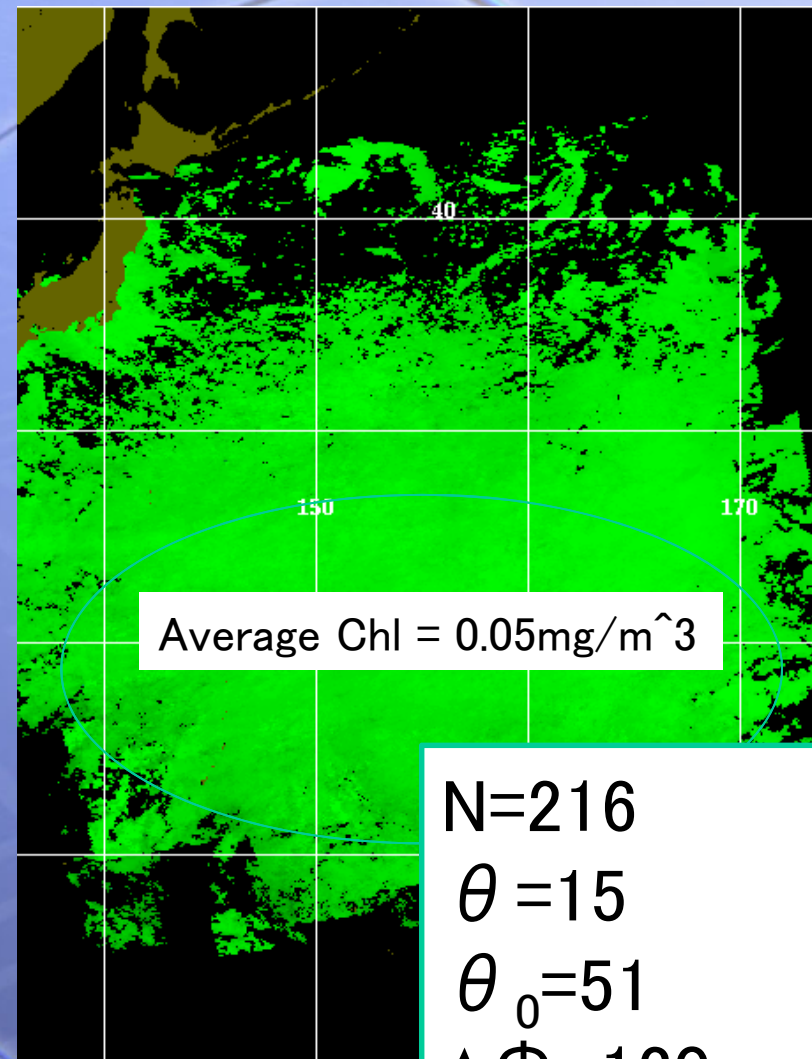
Aerosol type

mixture model

Tropospheric:Sea spray=40:60

Sampling Lt from MODIS/Aqua data

- Data in Jan. in 2013
- A2013003031000
- A2013005030000
- A2013007024500
- A2013008033000
- A2013009023500
- A2013010032000
- A2013015033500
- A2013016024000
- A2013017032500
- A2013019031000
- A2013021030000
- A2013024033000
- A2013027022500



$N=216$

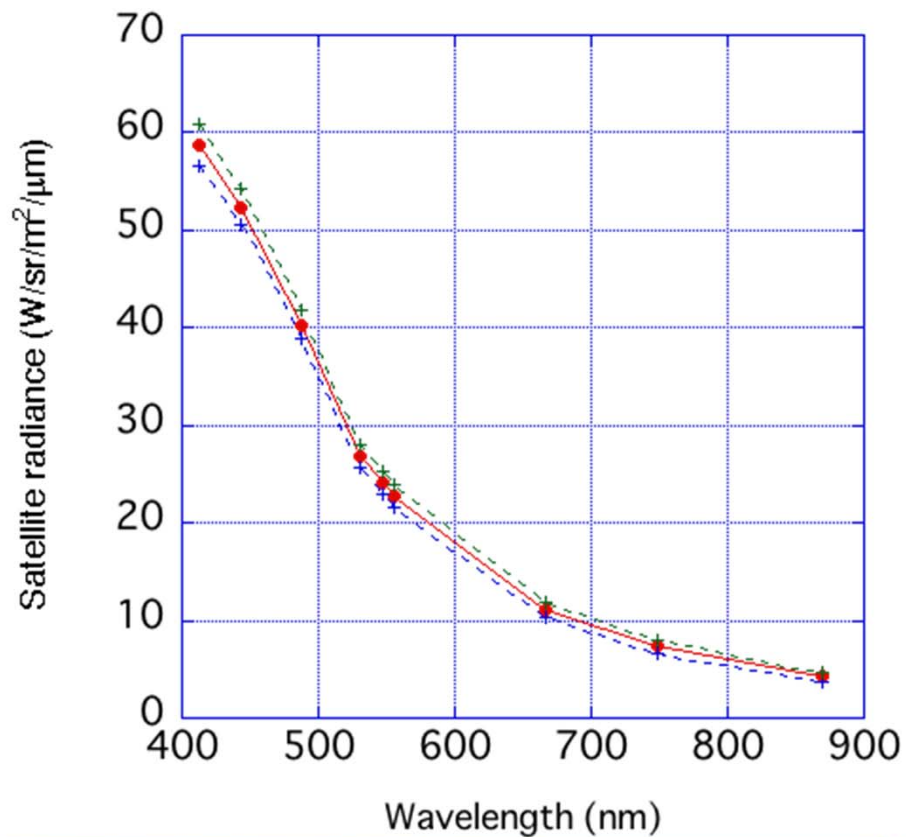
$\theta = 15$

$\theta_0 = 51$

$\Delta\Phi = 120$

$\tau_A(869) = 0.15$

Satellite radiance from MODIS/Aqua Jan. in 2013 on Pacific ocean



$N=216$

$\theta = 15$

$\theta_0 = 51$

$\Delta\Phi = 120$

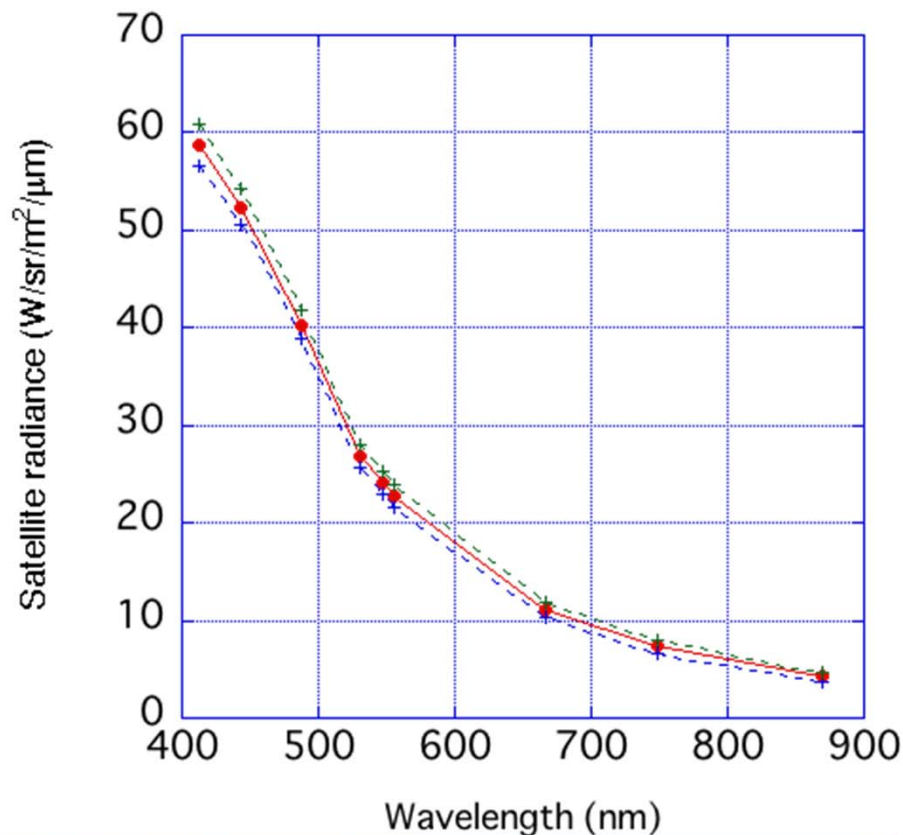
$\tau_A(869) = 0.15$

$L_t(869) : 4.25$

Std.dev.: 0.45

(W/(m²·sr·μm))

Satellite radiance from MODIS/Aqua Jan. in 2013 on Pacific ocean



$N=216$

$\theta = 15$

$\theta_0 = 51$

$\Delta\Phi = 120$

$\tau_A(865) = 0.15$

$Lt(869) : 4.25$

Std.dev.: 0.45

(W/(m²·sr·μm))

Simulation result $Lt(869) = 4.96$

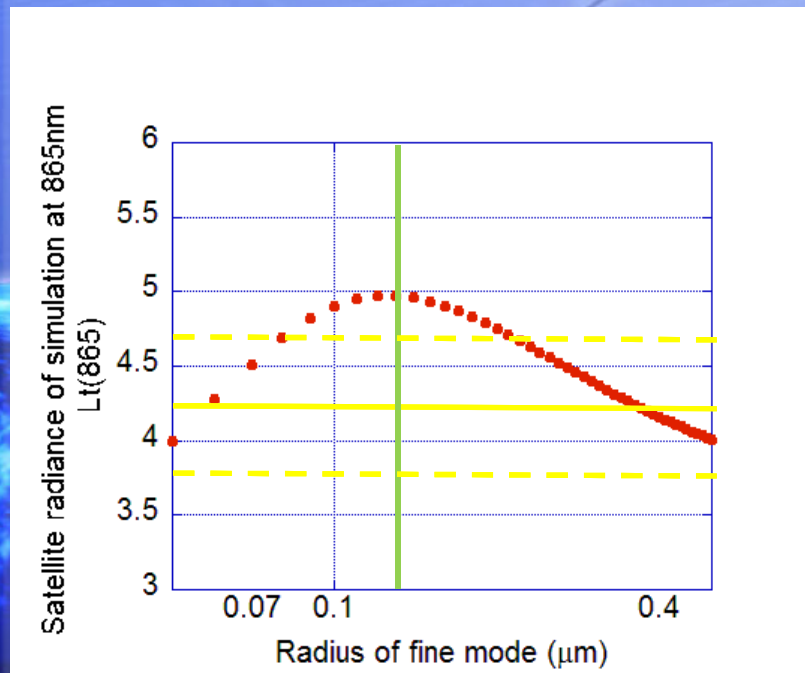
(W/(m²·sr·μm))

Aerosol model

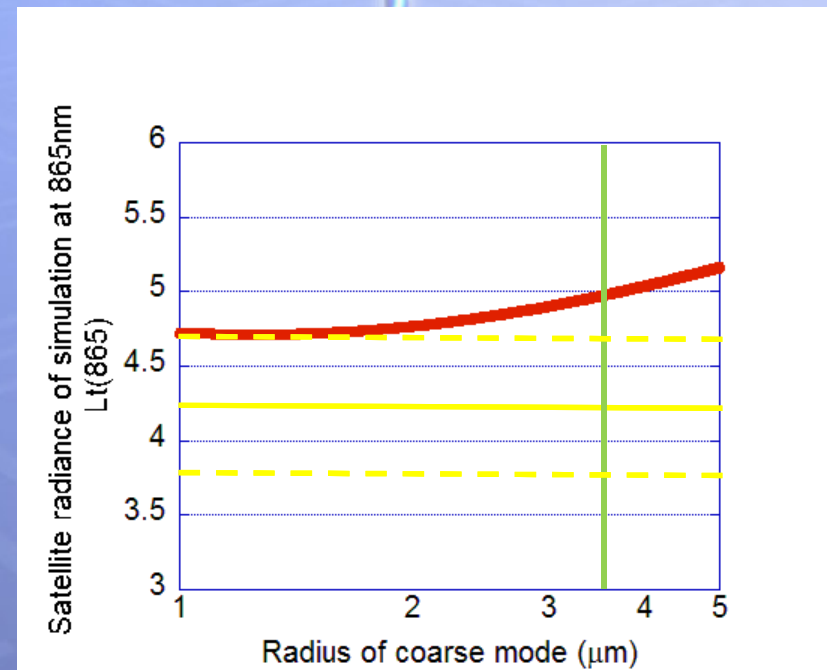
- Initial model was not applicable.
- Sensitivity test of Parameters
 - Small particle radius
 - Std.Dev. of small particle
 - Large particle radius
 - Std.Dev. of large particle
 - Ratio of volume size
 - Wind speed
 - Relative humidity
- Comparison Lt(865)

Result of sensitivity test

Radius of fine mode

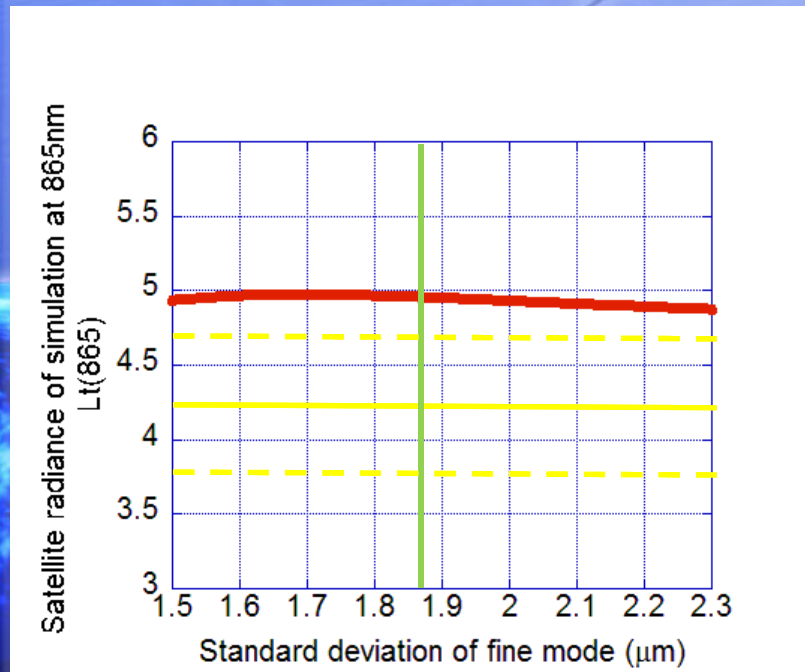


Radius of coarse mode

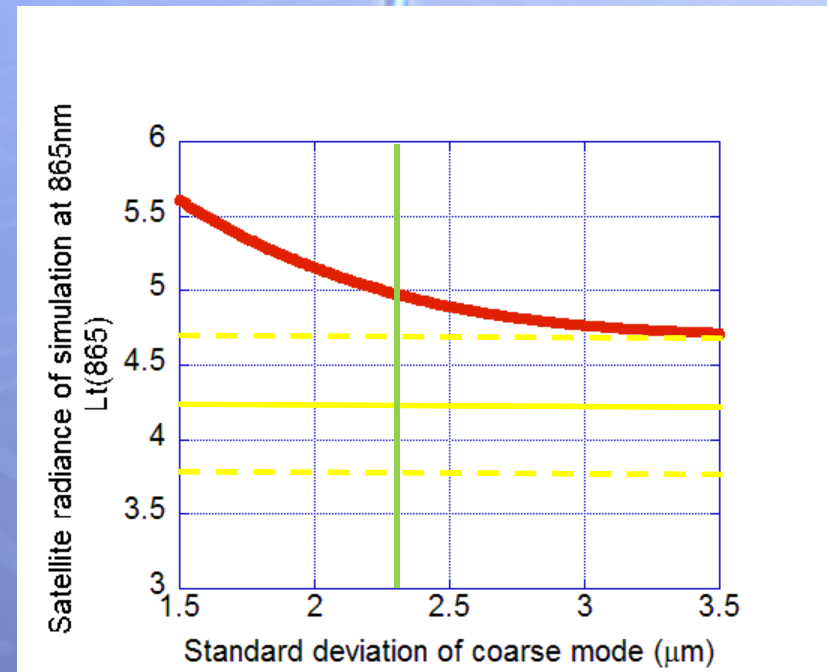


Result of sensitivity test

Std.dev. of fine mode



Std.dev. of coarse mode



Sensitivity test of aerosol parameters shows that average radius of fine mode is sensitive.

Cloud screening problems

- Highly turbid water

Manageable by
two-step masking

- may trip the cloud mask

- Sub-pixel cloud (incl. cloud edge)

- may affect atmospheric correction

- Cloud shadow

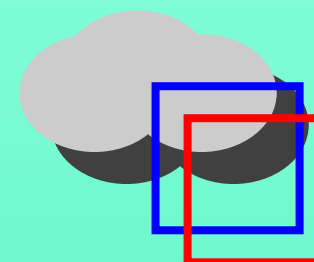
- “low Rrs555” mask may not enough

Bigger effect when

- inter-band registration is not perfect

- or

- inter-band time difference is large

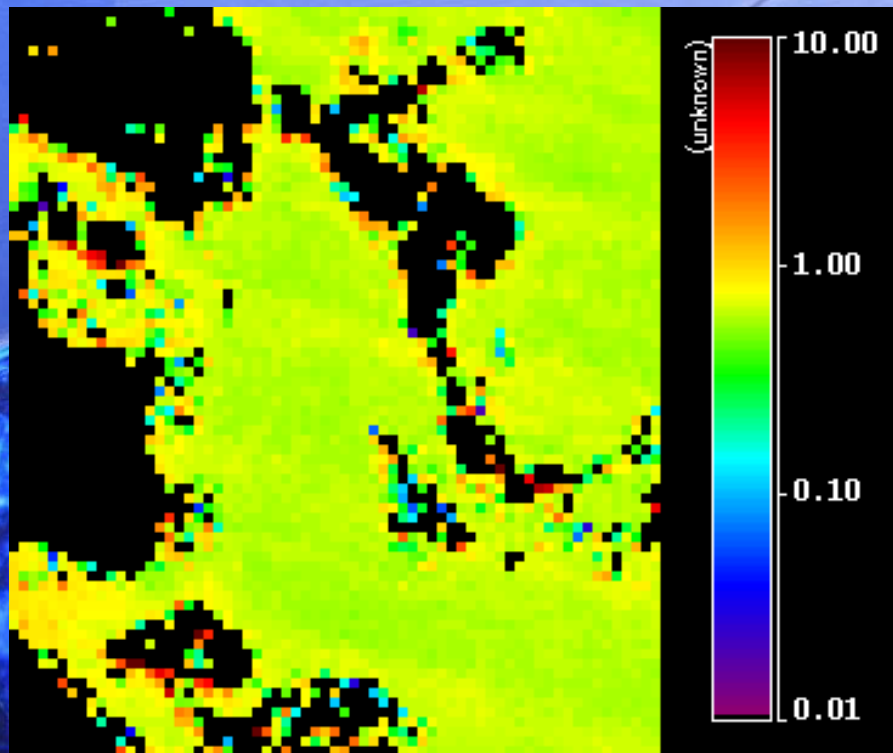


cause erroneous L2
pixel value

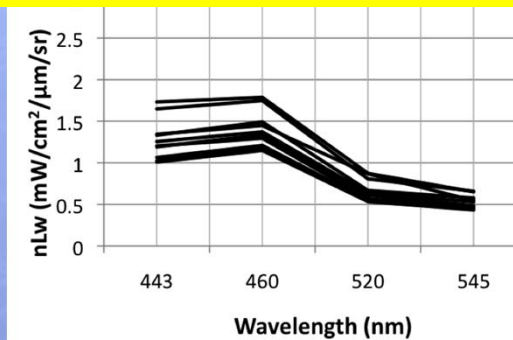
GLI, SGLI, ...

GOCI, ...

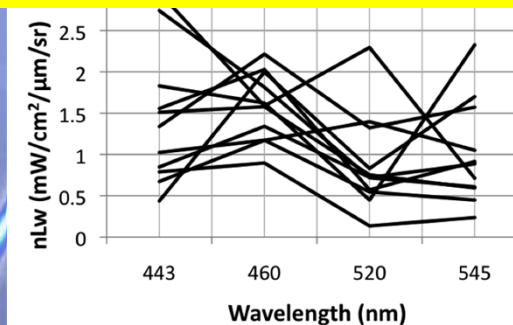
Anomalous features around cloud edge pixels (from GLI chl-a image of May 24, 2003, south of Japan)



nLw in “**Cloud-free**” area



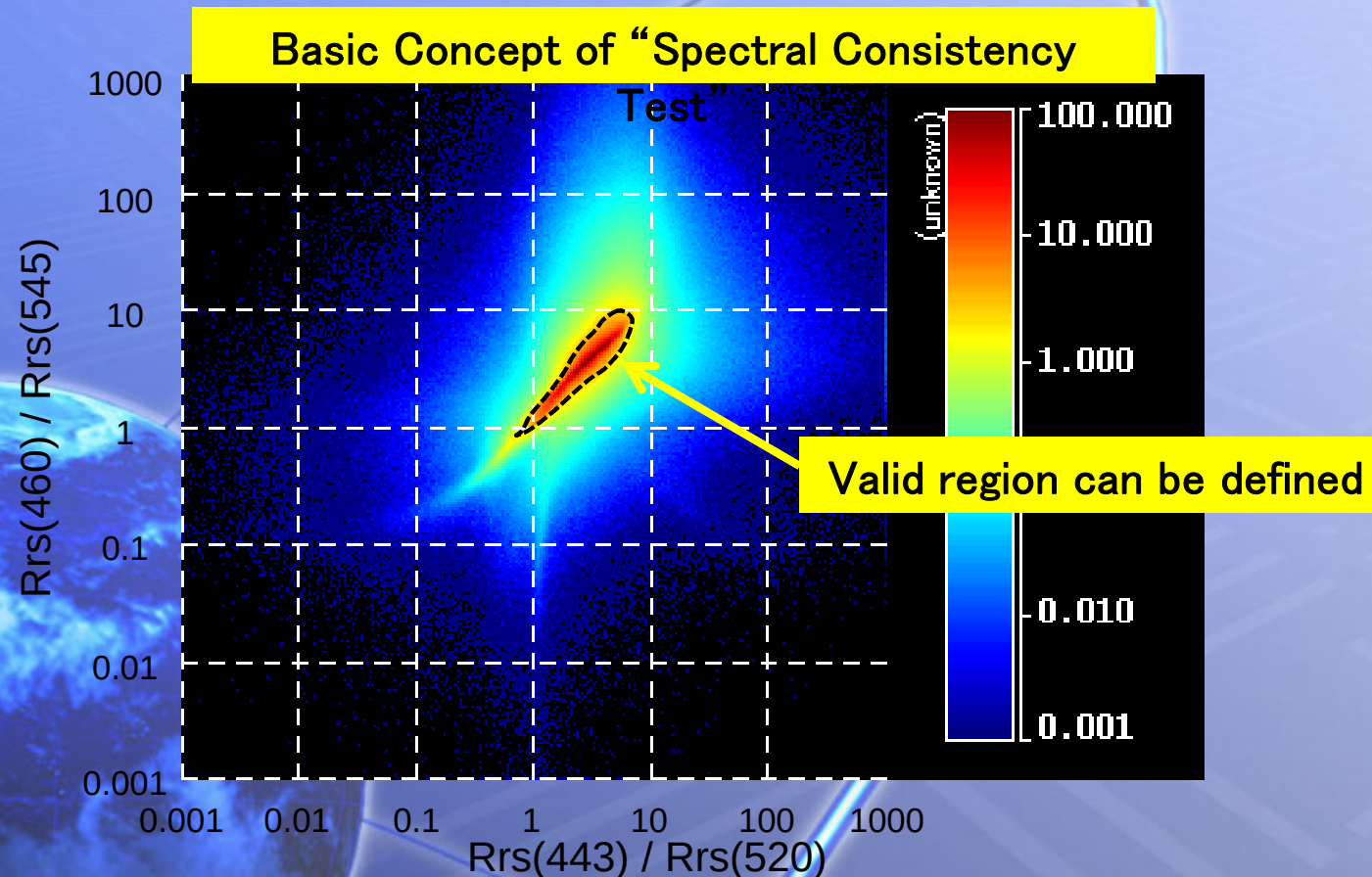
nLw in at “**Cloud-edge**”



Normalized water-leaving radiance
at 4 GLI bands.

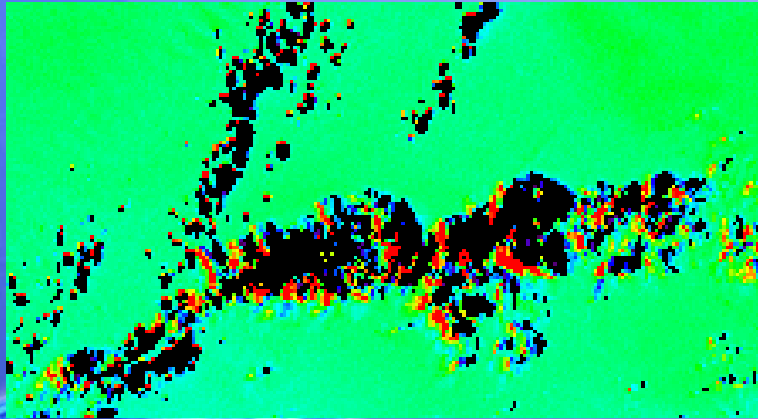
- a) Pixels sampled at cloud-free area, and
- b) sampled at cloud-adjacent pixels with
anomalous concentration estimate.

Inter-band Reflectance Consistency over a Weekly Global GLI Data

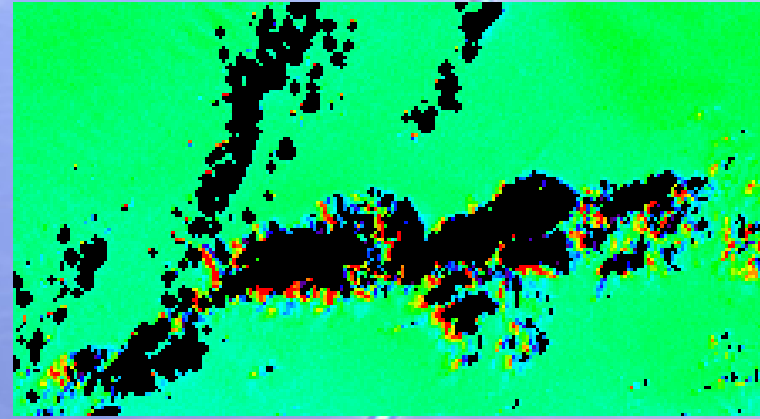


ADEOS-II / GLI
June 10-17, 2003
8day data set

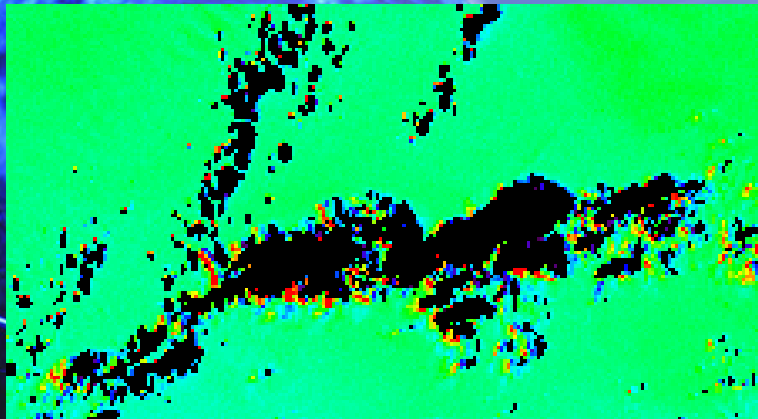
Sub-scene 1 (Typhoon case)



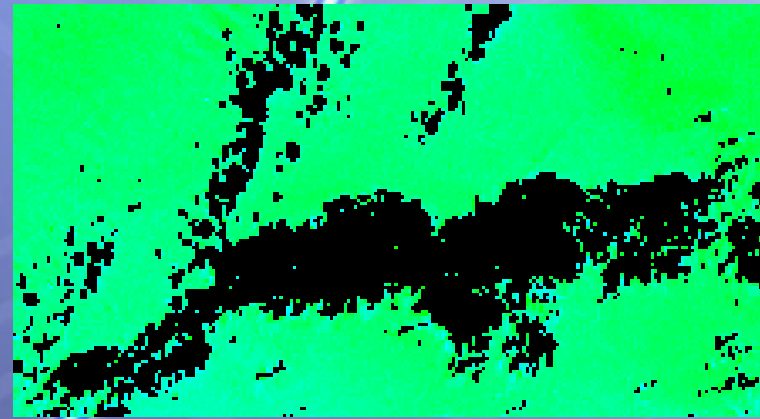
simple threshold



GOCI standard

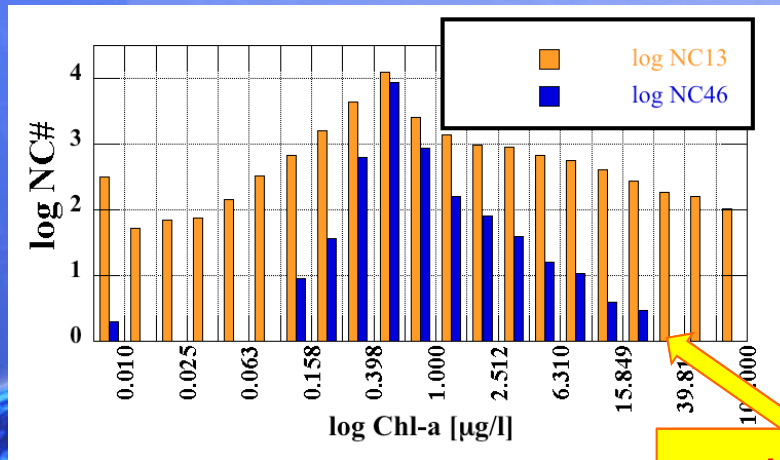


8-band threshold

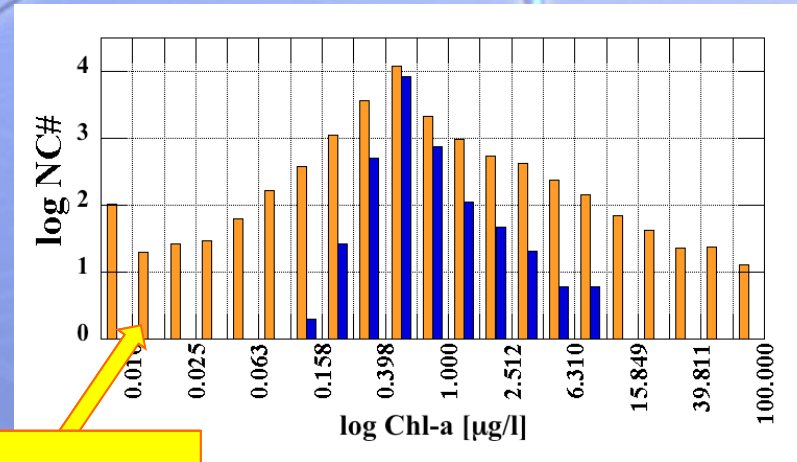


IRC

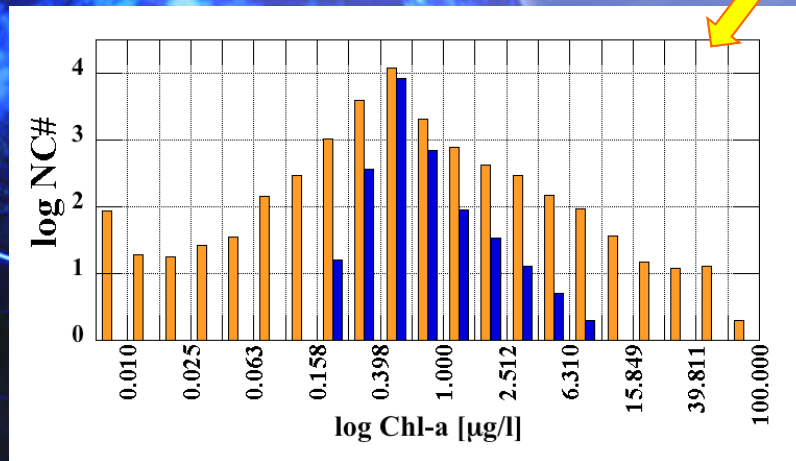
CSP Histogram for Typhoon Scene



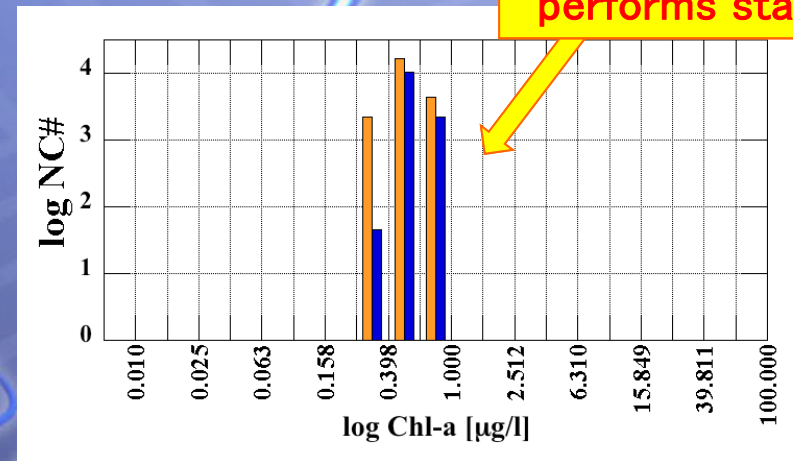
simple threshold



GOCI standard



8-band threshold



IRC

rather poor

performs stable

Summary

- Spectral dependency of aerosol optical thickness
 - relationship between Angstrom exponent aerosol type
 - Dust quantity related $\alpha(440,675)$ except for the ratio BC/SO_4 was higher (>0.2).
 - Requiring attention of component of fine mode, especially absorptive aerosol such as black carbon.



Summary

- Aerosol model to make look-up tables for aerosol reflectance correction
 - Proposed aerosol model was not applicable in this study.
 - Proposal model and sensitivity test of aerosol parameters
 - Sensitivity test of aerosol parameters shows that average radius of fine mode is sensitive.
- Inter-band Radiance Consistency method for cloud masking had good performance.

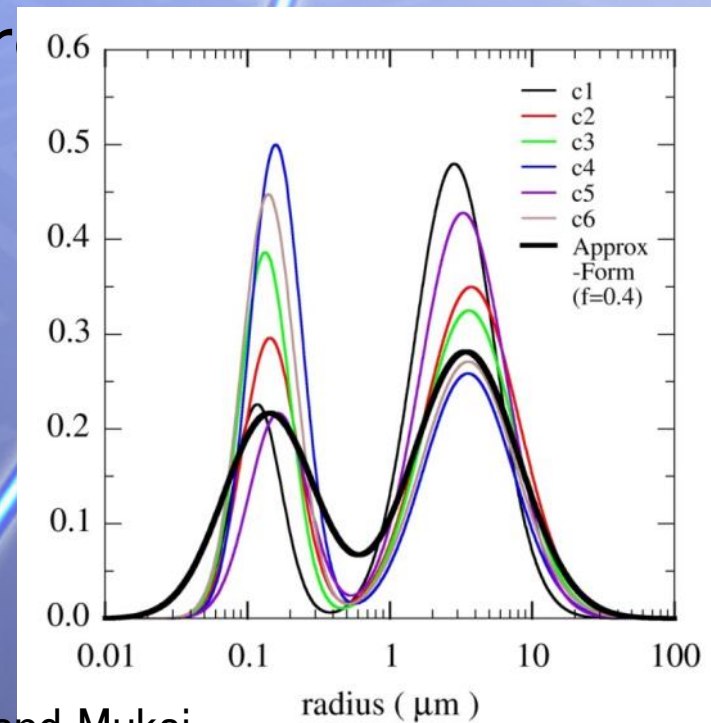
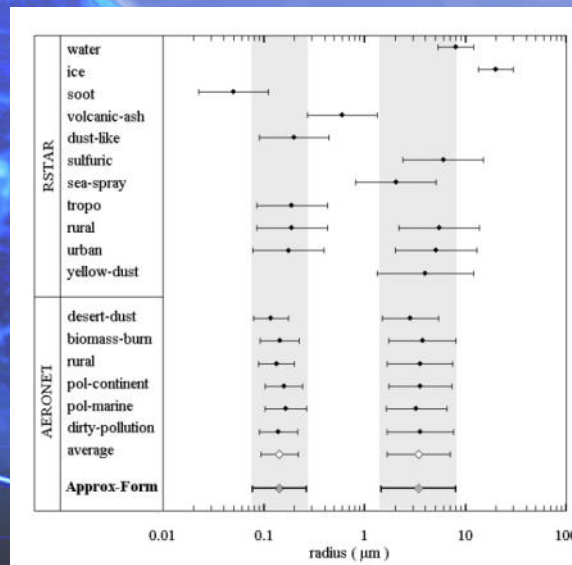


Backup slide



Aerosol correction

- Now we are making aerosol look-up table.
 - Aerosol model provided by atmosphere group
 - it derived from AERONET data
 - Rstar6b (Radiative transfer code)
- Absorptive aerosol corr



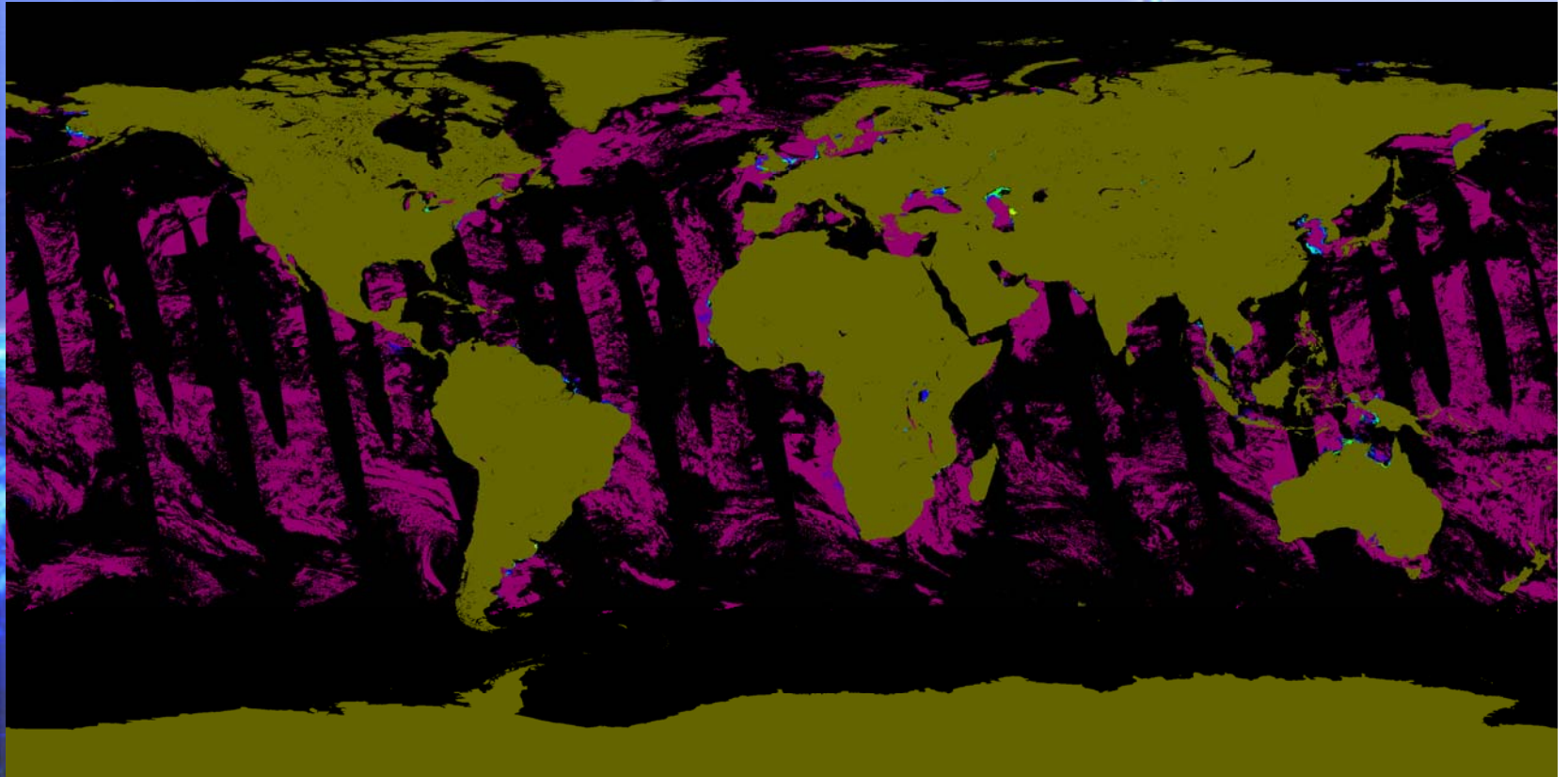
Provided by Nakata, Sano and Mukai
(Kinki university)

Atmospheric correction for high turbid area

- Detection of high turbid area
 - Threshold of total suspended matter concentration (TSM)
 - We developed simultaneously estimation algorithm of TSM concentration, AOT and Angstrom exponent using three bands at NIR.



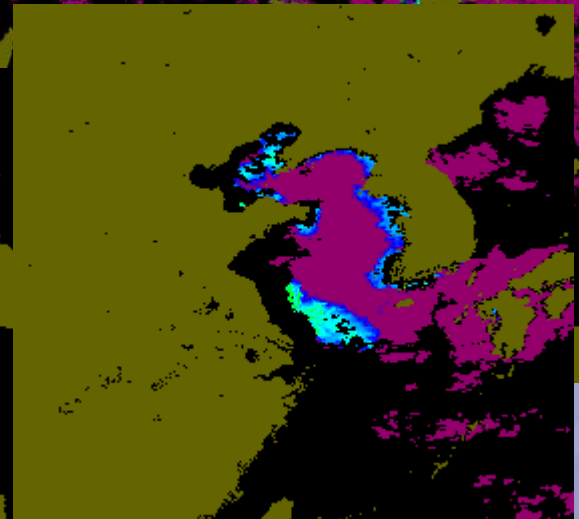
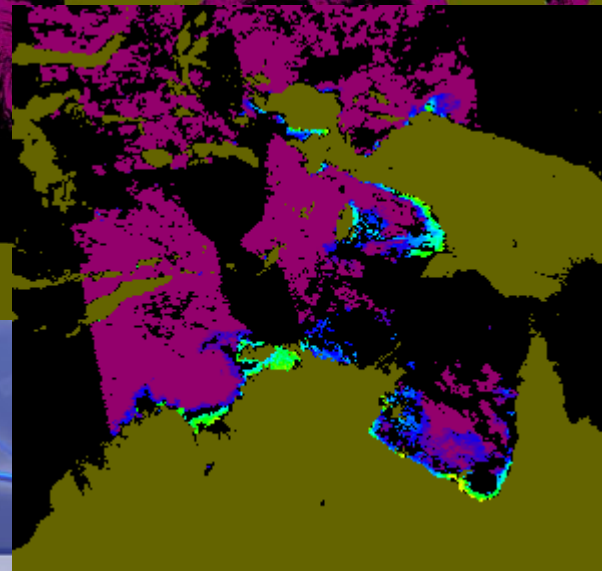
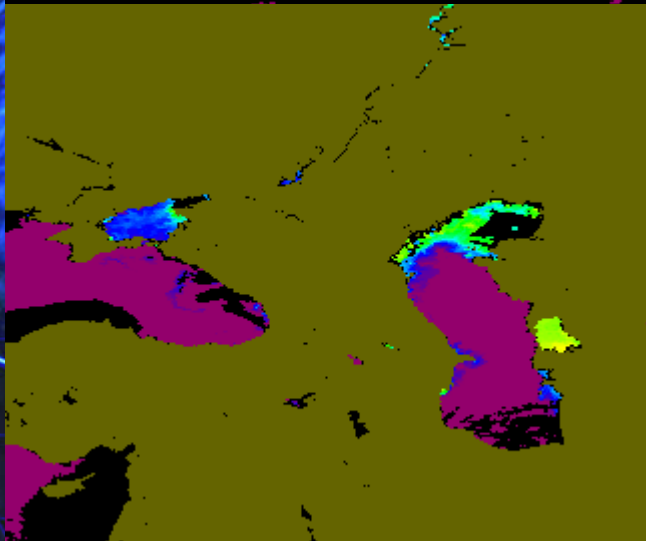
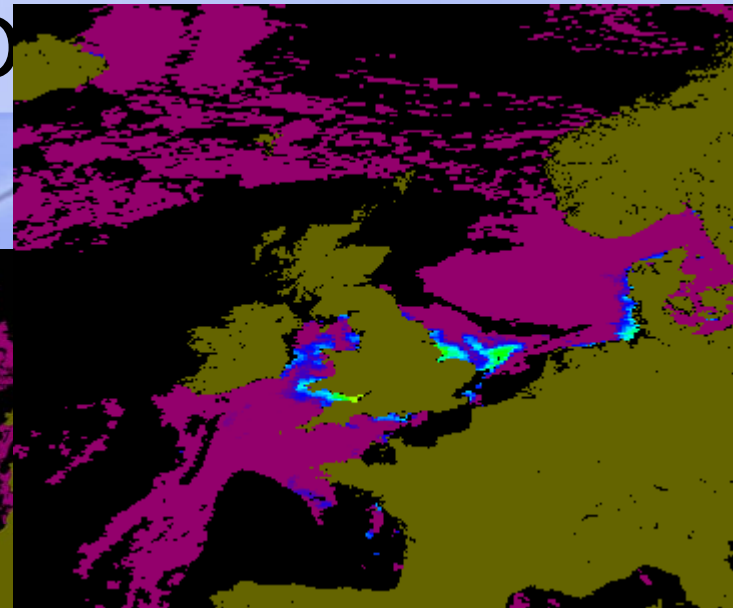
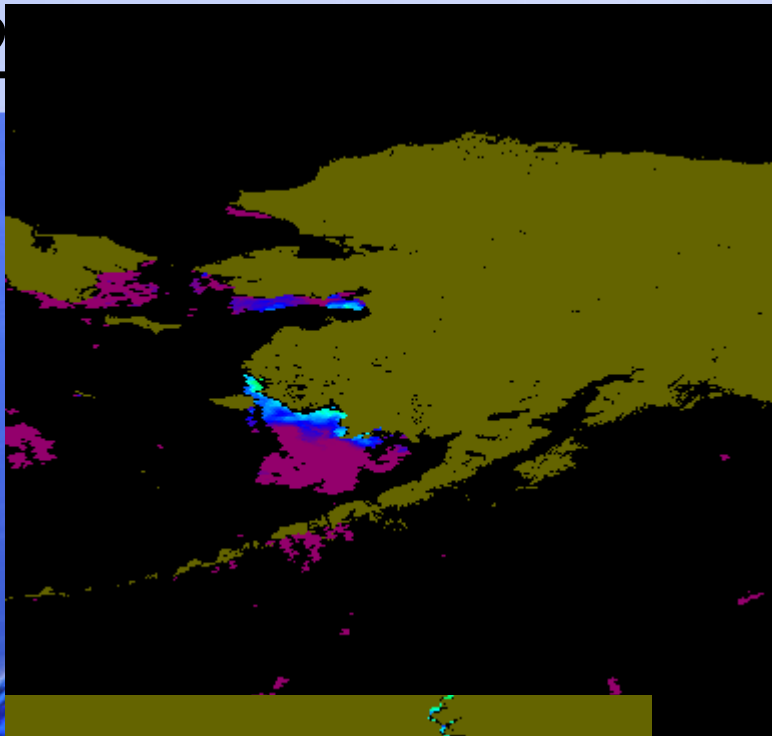
2013/5/3 TSM from MODIS three NIR bands



2

MODIS

S



Problem for SGLI

- SGLI has three NIR bands (670, 765, 865).
- It is difficult to use 765nm band because of oxygen absorption.
- Consider for Oxygen absorption correction
- Or use constant of angstrom exponent
- Laughly TSM estimation

